Blocking measures, hills, and hydrodynamics Joint with Jacob Calvert, Patrícia Goncalves and Katerina Michaelides

Márton Balázs

University of Bristol

Dynamics, random media and universality of complex physical systems Münster, 29 August, 2019.

Models

Asymmetric simple exclusion Zero range

Classical knowledge

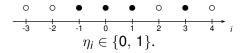
Asymmetric hydrodynamics Symmetric hydrodynamics

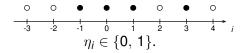
Blocking measures

ASEP ZRP Further models

Hills

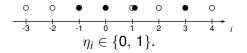
Microscopic model Hydrodynamics





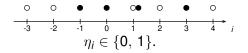
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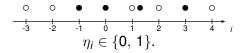
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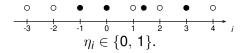
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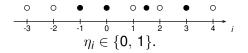
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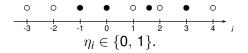
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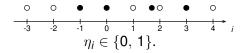
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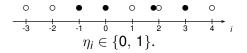
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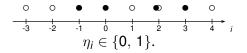
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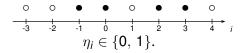
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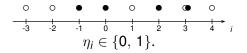
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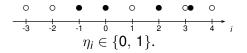
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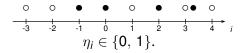
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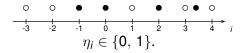
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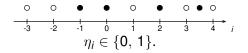
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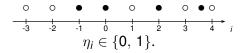
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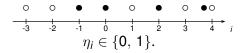
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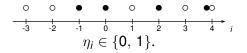
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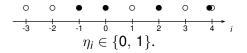
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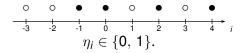
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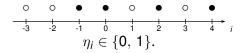
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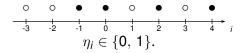
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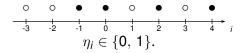
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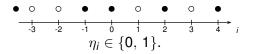
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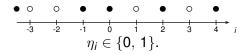
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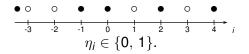
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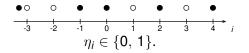
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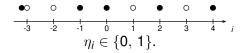
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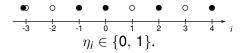
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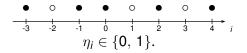
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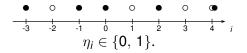
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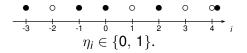
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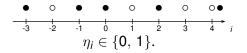
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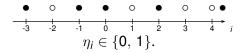
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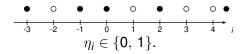
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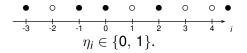
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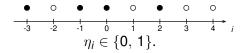
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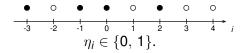
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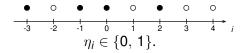
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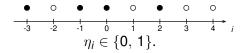
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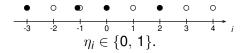
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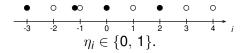
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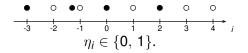
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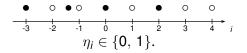
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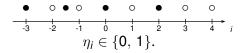
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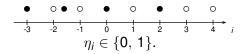
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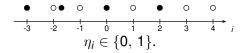
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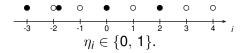
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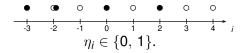
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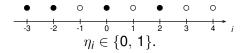
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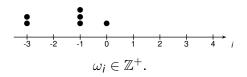
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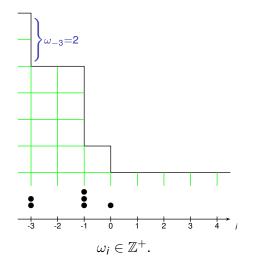
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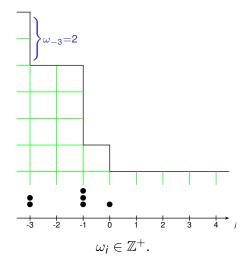


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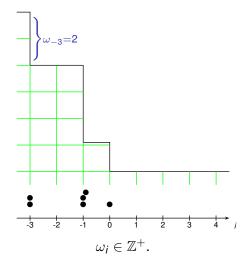
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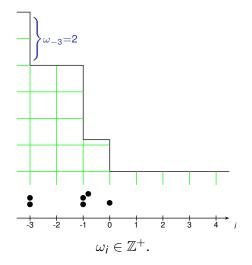




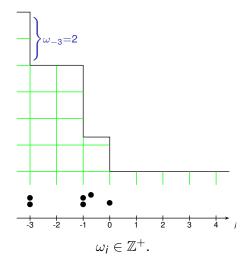
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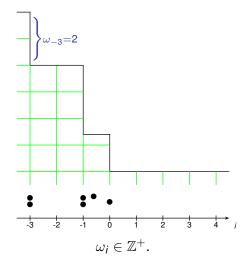
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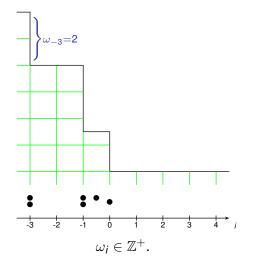
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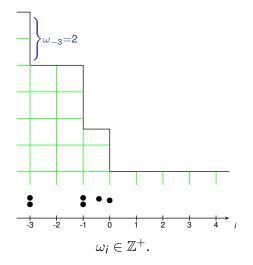
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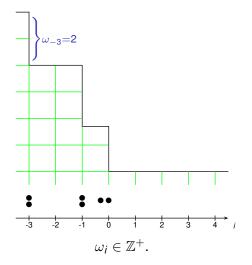
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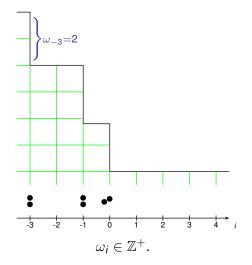
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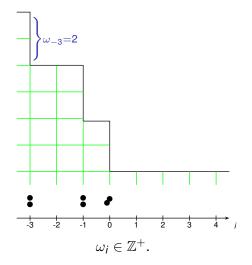
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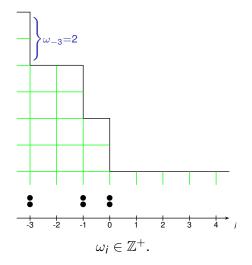
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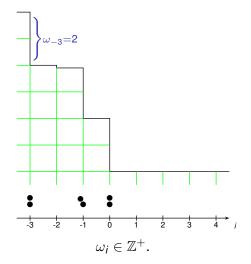
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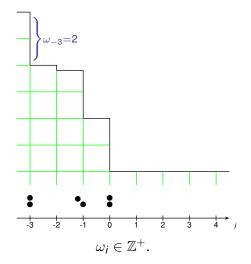
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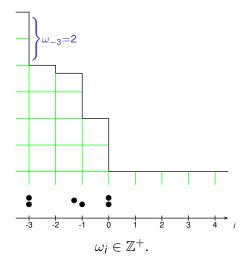
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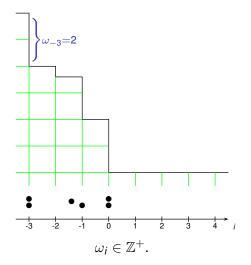
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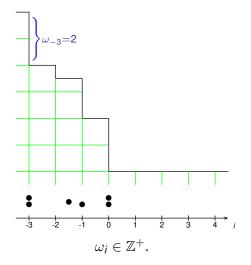
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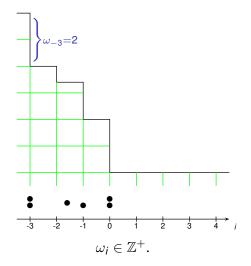
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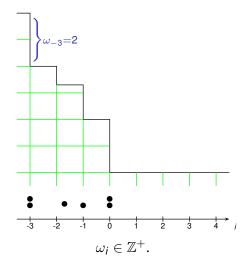
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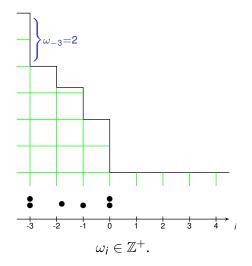
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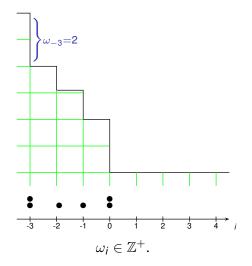
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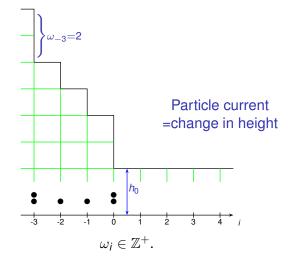
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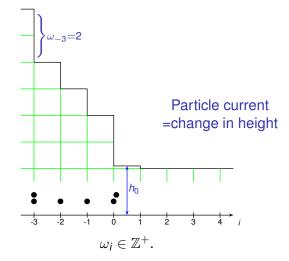
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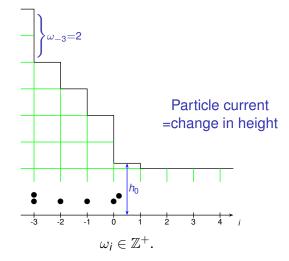
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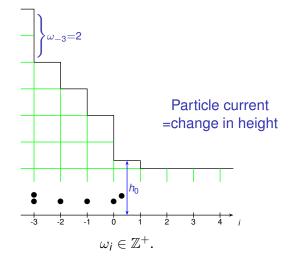
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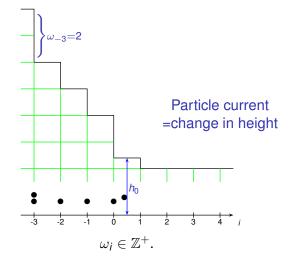
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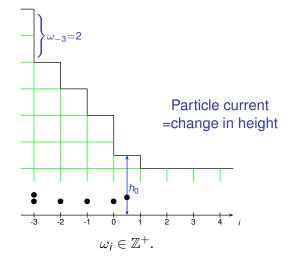
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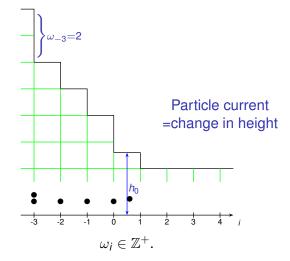
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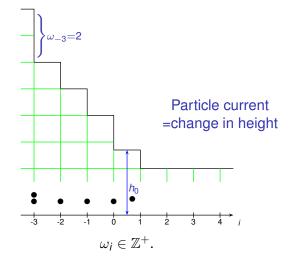
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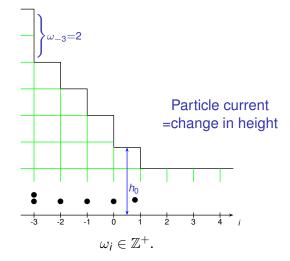
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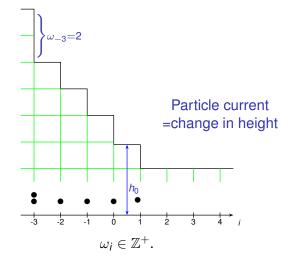
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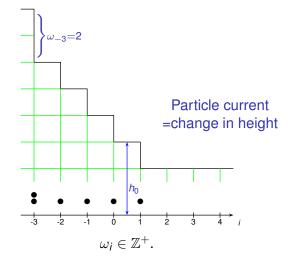
Particles jump



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Particles jump

We need *r* non-decreasing.

Examples:

- 'Classical' ZRP: $r(\omega_i) = \mathbf{1}\{\omega_i > \mathbf{0}\}.$
- Independent walkers: $r(\omega_i) = \omega_i$.

Translation invariant measures

No surprise that constant density is stationary:

<u>ASEP</u>: $\eta_i \sim \text{iid. Bernoulli}(\varrho)$.

<u>Classical ZRP:</u> $\omega_i \sim \text{iid. Geometric}(\frac{1}{1+a}).$

Independent walkers: $\omega_i \sim \text{iid. Poisson}(\varrho)$.

These are the only extremal translation-invariant distributions.

Take p > q = 1 - p, and AZRP with right rate $p \cdot r(\omega_i)$, left rate $q \cdot r(\omega_i)$.

$$\frac{\mathsf{d}}{\mathsf{d}\tau} \mathbf{E}\omega_i = p \mathbf{E}r(\omega_{i-1}) + q \mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_i)$$

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Ballistic scaling (zoom out and speed up by factor *L*):

- $\blacktriangleright \ \varrho(t, x) = \mathbf{E}\omega_{Lx}(Lt);$
- ► also define $G(\varrho) = \mathbf{E}^{\varrho} r(\omega)$: $\frac{d}{d(\tau/L)} \mathbf{E}\omega_i = qL(\mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_i)) - pL(\mathbf{E}r(\omega_i) - \mathbf{E}r(\omega_{i-1}))$

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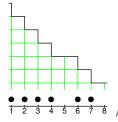
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$$rac{\partial}{\partial t}arrho(t,\,x)+(
ho-q)rac{\partial}{\partial x}G(arrho(t,\,x))=0$$

<u>Classical ZRP</u>: $G(\varrho) = \mathbf{E}^{\varrho} r(\omega) = \mathbf{E}^{\varrho} \mathbf{1} \{ \omega > 0 \} = \frac{\varrho}{1+\varrho}$ concave, Burgers-type equation.

Independent walkers: $G(\varrho) = \mathbf{E}^{\varrho} r(\omega) = \mathbf{E}^{\varrho} \omega = \varrho$ linear, transport equation.



The stationary solution is constant density, linear slope.

Take $p = q = \frac{1}{2}$, and SZRP with right rate $\frac{1}{2}r(\omega_i)$, left rate $\frac{1}{2}r(\omega_i)$.

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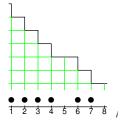
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The stationary solution is constant density, linear slope, or linearly changing G with current.

Hills

Can we model sedimentation and erosion processes with these surfaces?

Issues:

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► Hills are not always straight ↔ translation invariance.

Hills

Can we model sedimentation and erosion processes with these surfaces?

Issues:

- Hills are not always straight \leftrightarrow translation invariance.
- ▶ Most hillslopes are rather stationary ↔ particle current.

Convex hills



Wikipedia

Concave hills



Stockphotos4free

Product blocking measures

Solution: block particles (no current) and make their rates asymmetric (non-constant density).

Can we have a reversible stationary distribution in product form:

$$\underline{\mu}(\underline{\omega}) = \bigotimes_{i} \mu_{i}(\omega_{i});$$

$$\underline{\mu}(\underline{\omega}) \cdot \mathsf{rate}(\underline{\omega} \to \underline{\omega}^{i \frown i+1}) = \underline{\mu}(\underline{\omega}^{i \frown i+1}) \cdot \mathsf{rate}(\underline{\omega}^{i \frown i+1} \to \underline{\omega}) \quad ?$$

Here

$$\underline{\omega}^{i \frown i+1} = \underline{\omega} - \underline{\delta}_i + \underline{\delta}_{i+1}.$$

Asymmetric simple exclusion

 $\mu(\eta) \cdot \mathsf{rate}(\eta \to \eta^{i \frown i+1}) = \mu(\eta^{i \frown i+1}) \cdot \mathsf{rate}(\eta^{i \frown i+1} \to \eta)$ <u>ASEP:</u> $\mu_i \sim \text{Bernoulli}(\varrho_i); \quad \neg \eta$ $\rho_i(1 - \rho_{i+1}) \cdot \rho = (1 - \rho_i)\rho_{i+1} \cdot q$ Solution: $\varrho_i = \frac{(\frac{\rho}{q})^{i-c}}{1 + (\frac{\rho}{2})^{i-c}} = \frac{1}{(\frac{q}{2})^{i-c} + 1}$ 0 С

Asymmetric simple exclusion

$$\underline{\mu}(\underline{\omega}) \cdot \mathsf{rate}(\underline{\omega} \to \underline{\omega}^{i \frown i+1}) = \underline{\mu}(\underline{\omega}^{i \frown i+1}) \cdot \mathsf{rate}(\underline{\omega}^{i \frown i+1} \to \underline{\omega}) \quad ?$$

AZRP, classical:

 $\mu_i(\omega_i)\mu_{i+1}(\omega_{i+1})\cdot p\mathbf{1}\{\omega_i>0\}=\mu_i(\omega_i-1)\mu_{i+1}(\omega_{i+1}+1)\cdot q$

Solution:
$$\mu_i \sim \text{Geometric}\Big(1 - \Big(\frac{p}{q}\Big)^{i-\text{const}}\Big).$$

$$\underline{\mu}(\underline{\omega}) \cdot \mathsf{rate}(\underline{\omega} \to \underline{\omega}^{i \frown i+1}) = \underline{\mu}(\underline{\omega}^{i \frown i+1}) \cdot \mathsf{rate}(\underline{\omega}^{i \frown i+1} \to \underline{\omega}) \quad ?$$

AZRP, independent walkers:

$$\mu_i(\omega_i)\mu_{i+1}(\omega_{i+1})\cdot p\omega_i = \mu_i(\omega_i-1)\mu_{i+1}(\omega_{i+1}+1)\cdot q(\omega_{i+1}+1)$$

Solution:
$$\mu_i \sim \text{Poisson}\left(\left(\frac{p}{q}\right)^{i-\text{const}}\right)$$
.



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Other models can be stood up:

- ASEP
- q-exclusion
- Katz-Lebowitz-Spohn model

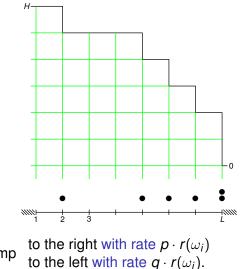
Product blocking measures

They are also very handy, due to reversibility.

Take a stationary, reversible Markov chain. Cut any of its edges. It stays reversible stationary w.r.t. the same distribution.

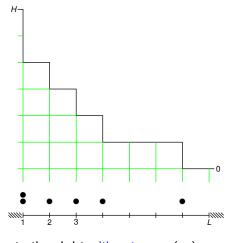
In our case: freeze the boundaries to obtain a stationary hill slope.

Our choice: AZRP with frozen boundaries. p > q: convex



Particles jump

Our choice: AZRP with frozen boundaries. p < q: concave



Particles jump

to the right with rate $p \cdot r(\omega_i)$ to the left with rate $q \cdot r(\omega_i)$.

Notice:

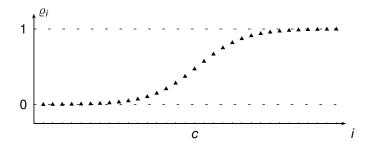
- The height of the hill H is conserved, the product measure is not ergodic.
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We won't be bothered by this.

Work in progress...

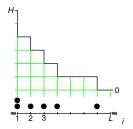


A blocking measure is a microscopic object. Here is its scaling

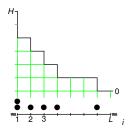
limit:
$$\xrightarrow{e^{(x)}}_{x}$$
, not very interesting.

Micro Hydro

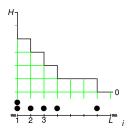
Hydrodynamics



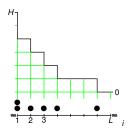
Scaling parameter: L



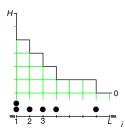
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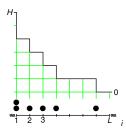


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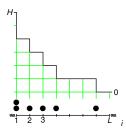


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For AZRP (rates $p \cdot r(\omega_i)$ right and $q \cdot r(\omega_i)$ left):

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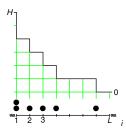


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which dictates diffusive scaling:

$$\blacktriangleright \ p = \frac{1}{2} + \frac{\gamma}{L}, \ q = \frac{1}{2} - \frac{\gamma}{L};$$

$$\blacktriangleright \ \varrho(t, x) = \mathbf{E}\omega_{Lx}(L^2t);$$

• also define $G(\varrho) = \mathbf{E}^{\varrho} r(\omega)$:

$$\begin{aligned} \frac{\mathsf{d}}{\mathsf{d}(\tau/L^2)} \mathbf{E}\omega_i &= \frac{L^2}{2} \big(\mathbf{E}r(\omega_{i-1}) - 2\mathbf{E}r(\omega_i) + \mathbf{E}r(\omega_{i+1}) \big) \\ &- \gamma L \big(\mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_{i-1}) \big), \\ \frac{\partial}{\partial t} \varrho(t, x) &= \frac{1}{2} \frac{\partial^2}{\partial x^2} G(\varrho(t, x)) - 2\gamma \frac{\partial}{\partial x} G(\varrho(t, x)), \quad (0 < x < 1). \end{aligned}$$

$$p = \frac{1}{2} + \frac{\gamma}{L} , \ q = \frac{1}{2} - \frac{\gamma}{L} ;$$
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Also:
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$$\frac{\partial}{\partial t}\varrho(t, x) = \frac{1}{2}\frac{\partial^2}{\partial x^2}G(\varrho(t, x)) - 2\gamma \frac{\partial}{\partial x}G(\varrho(t, x)) \quad (0 < x < 1)$$

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Convection-diffusion type equation with Robin boundary.

$$\frac{\partial}{\partial t}\varrho(t, x) = \frac{1}{2}\frac{\partial^2}{\partial x^2}G(\varrho(t, x)) - 2\gamma\frac{\partial}{\partial x}G(\varrho(t, x)) \quad (0 < x < 1)$$
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Convection-diffusion type equation with Robin boundary.

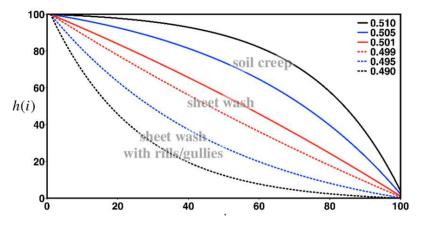
Doing the proper derivation is work in progress.

The time-stationary solution $G(\varrho(x)) = Ce^{4\gamma x}$ is consistent with the stationary blocking measure.

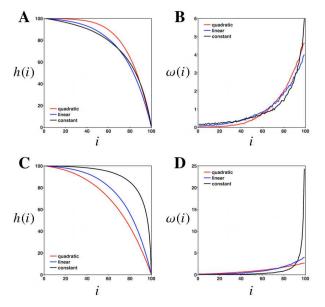
Micro Hydro

The stationary slope

 $G(\varrho(x)) = C \mathrm{e}^{4\gamma x}$



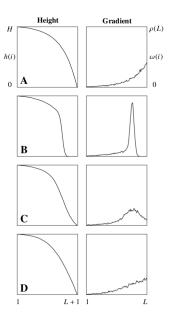
The stationary slope

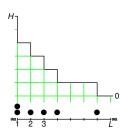


Space scale: $x \in [0, 1] \Leftrightarrow we \in hill$.

Problem 1: The stationary hillslope will not tell us the time scale.

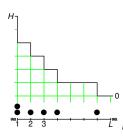
~ Observe relaxation to stationarity in Nature and in the PDE.





Problem 2: Geologists want a prediction for the *hill particle flux*, and the *distance travelled by hill particles*.

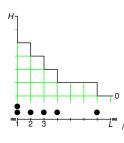
Notice: Hill particles \neq our particles.



Problem 2: Geologists want a prediction for the *hill particle flux*, and the *distance travelled by hill particles*.

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This is not part of the core argument, instead, is done by heuristics:

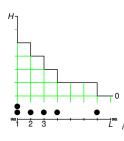


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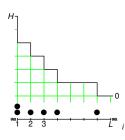


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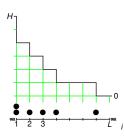


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- Average hill particle flux is the same across the hill (reversibility), but this is not provided by the model.



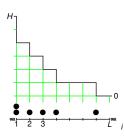
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One can then give an expected distance travelled by a hill particle. *Thank you.*