

Blocking measures, hills, and hydrodynamics

Joint with

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Márton Balázs

University of Bristol

Probability and NonLocal PDEs
Swansea, 17 September, 2018.

Models

- Asymmetric simple exclusion
- Zero range

Classical knowledge

- Asymmetric hydrodynamics
- Symmetric hydrodynamics

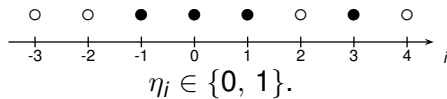
Blocking measures

- ASEP
- ZRP
- Further models

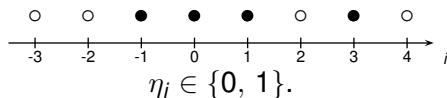
Hills

- Microscopic model
- Hydrodynamics

Asymmetric simple exclusion



Asymmetric simple exclusion



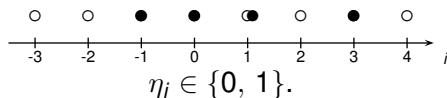
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Asymmetric simple exclusion



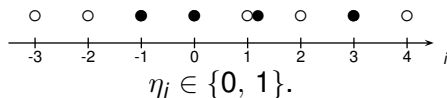
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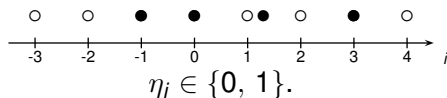
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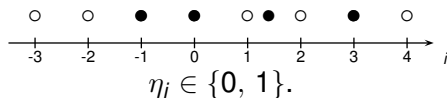
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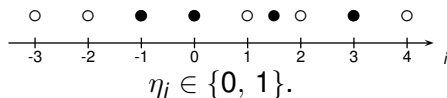
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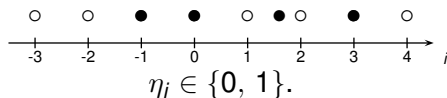
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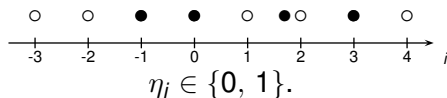
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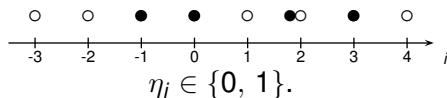
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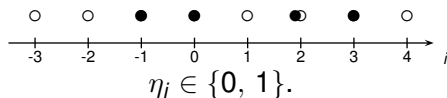
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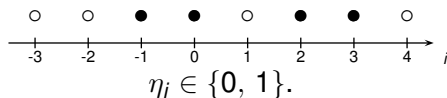
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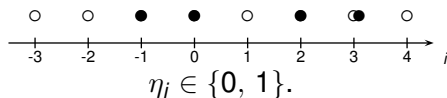
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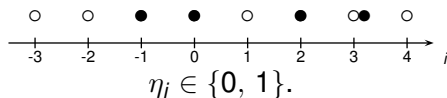
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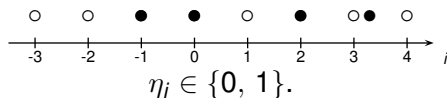
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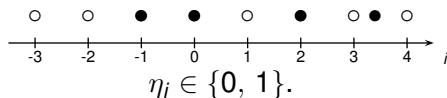
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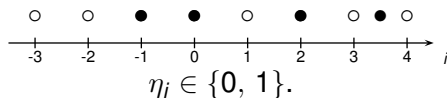
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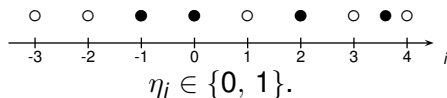
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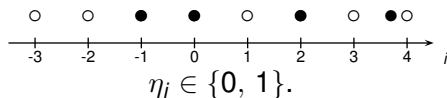
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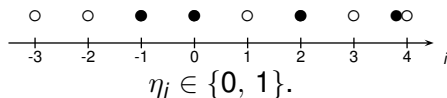
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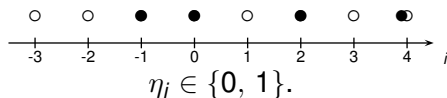
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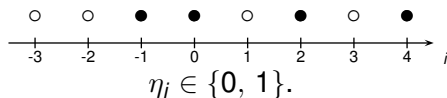
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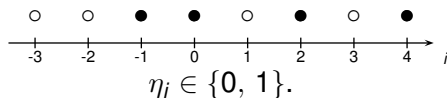
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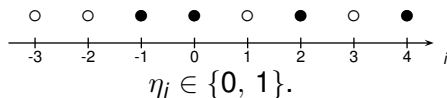
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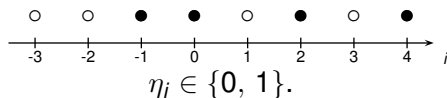
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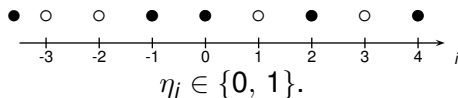
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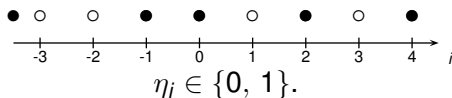
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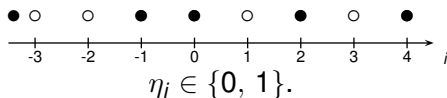
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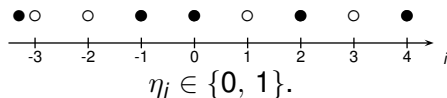
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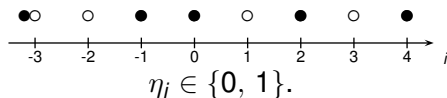
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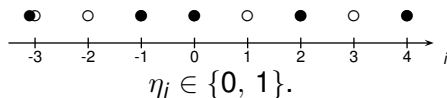
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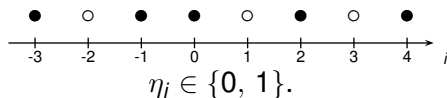
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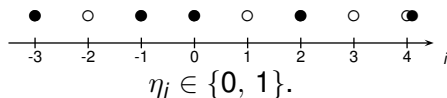
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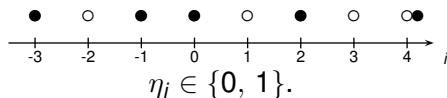
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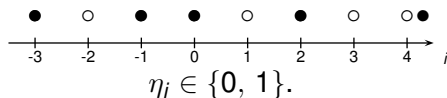
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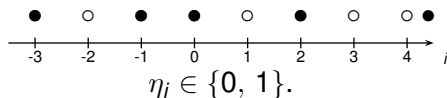
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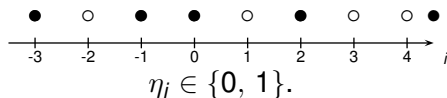
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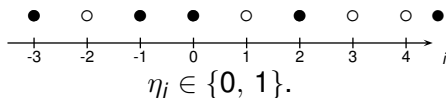
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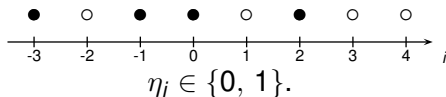
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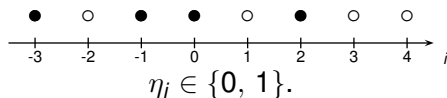
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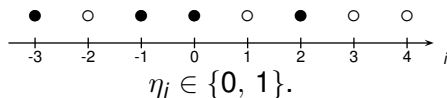
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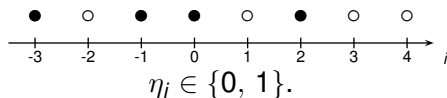
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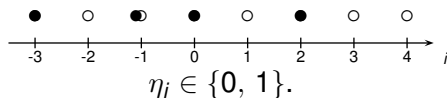
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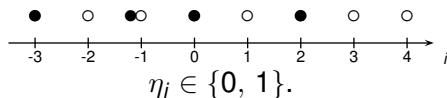
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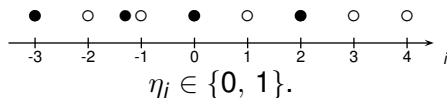
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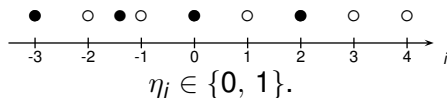
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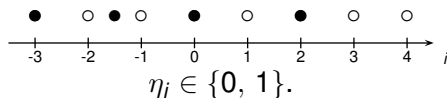
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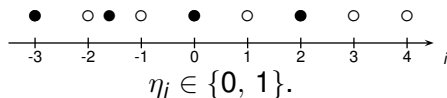
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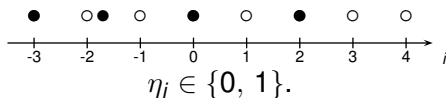
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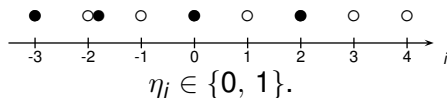
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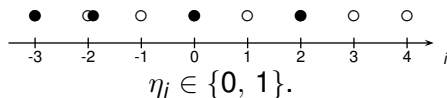
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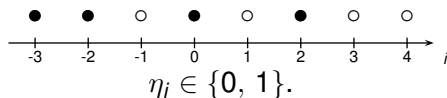
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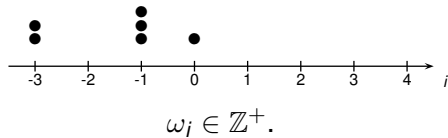
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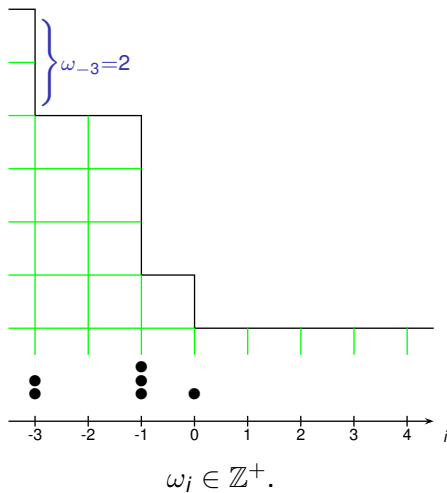
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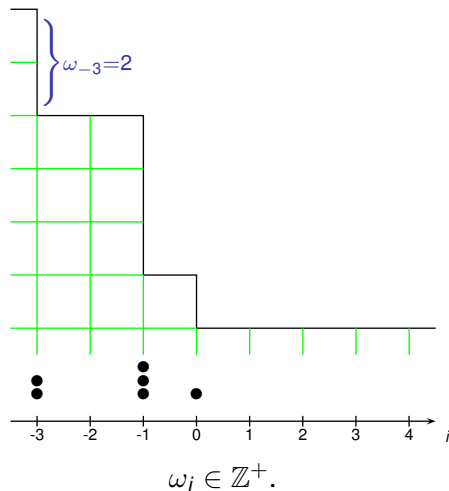
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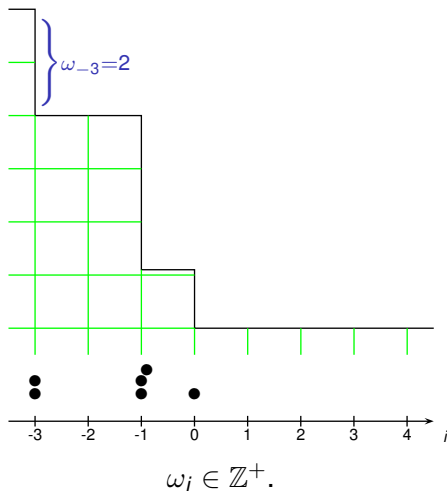


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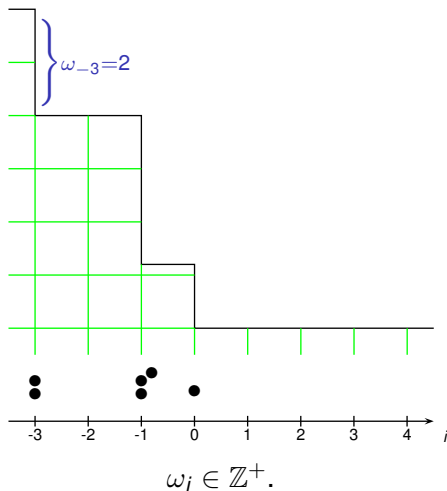
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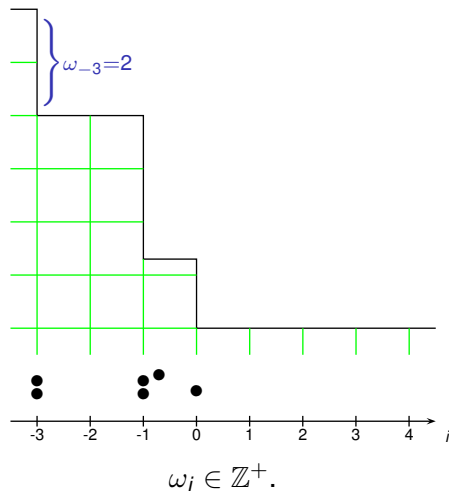
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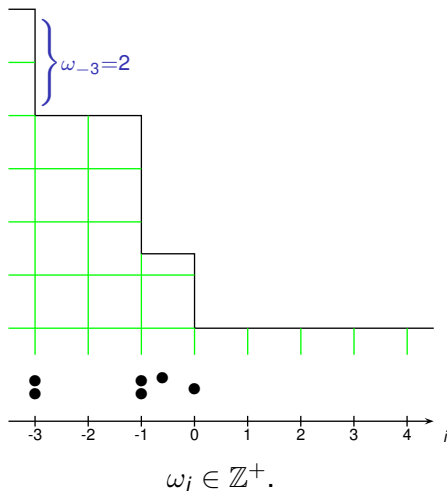
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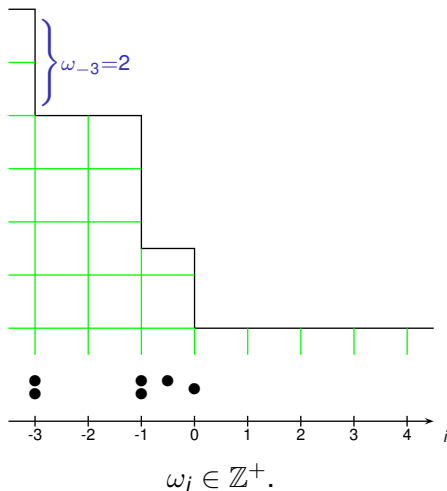
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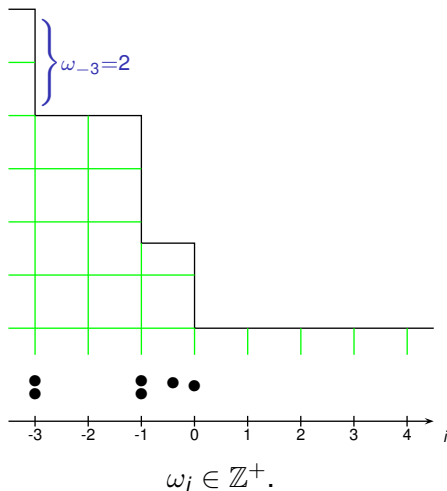
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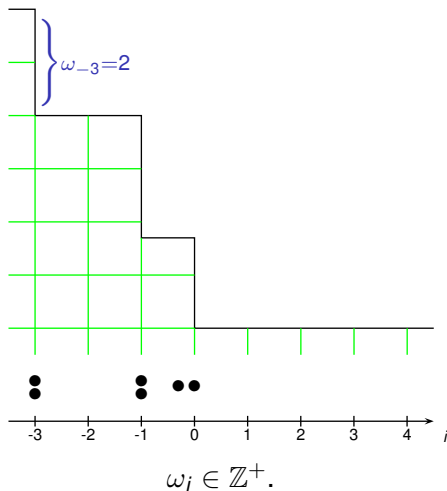
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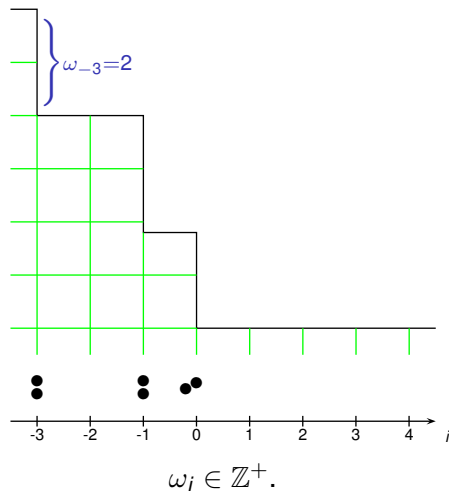
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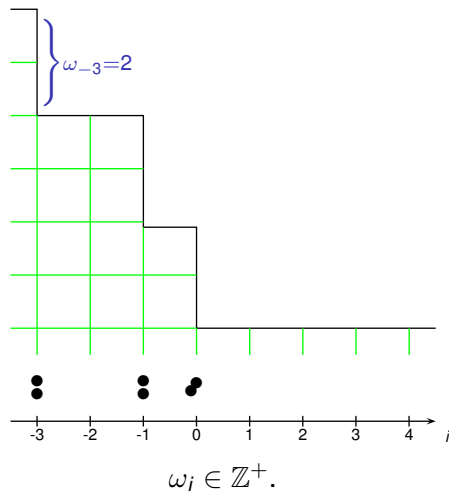
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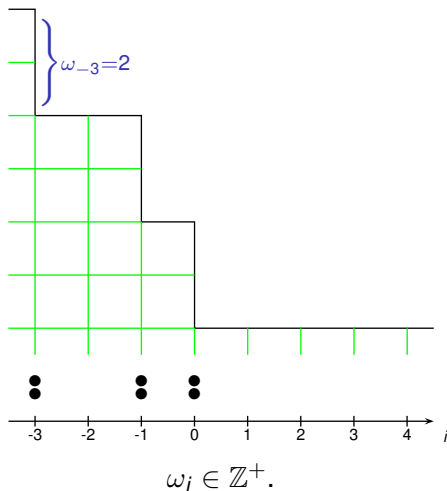
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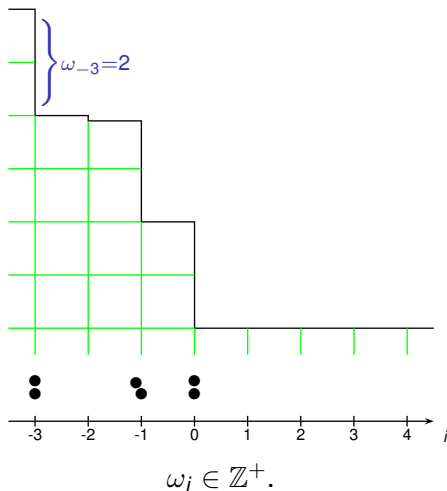
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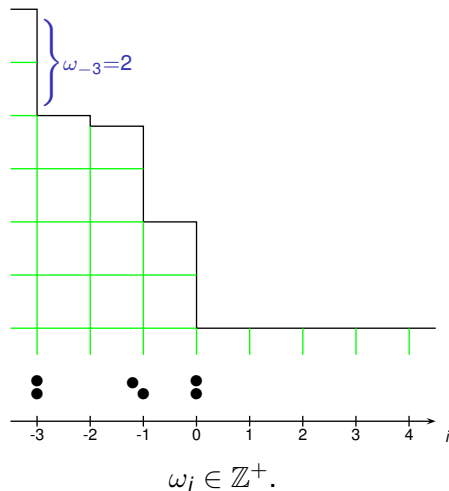
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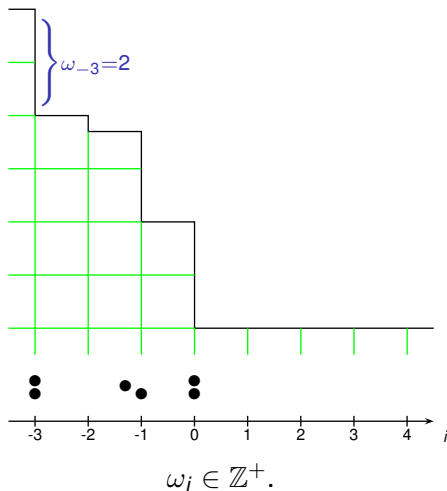
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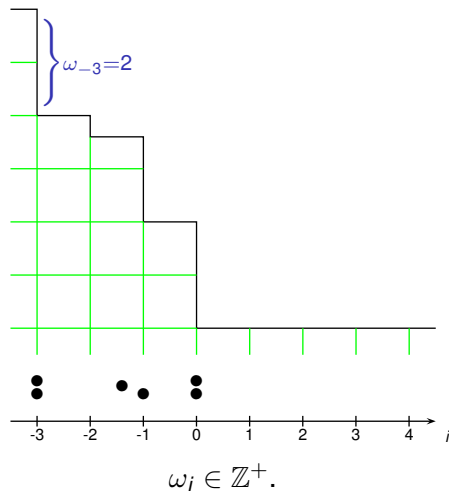
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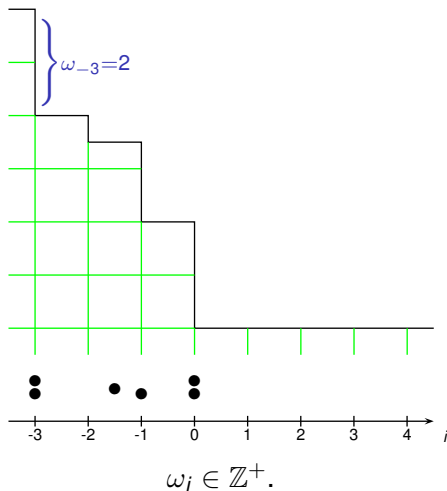
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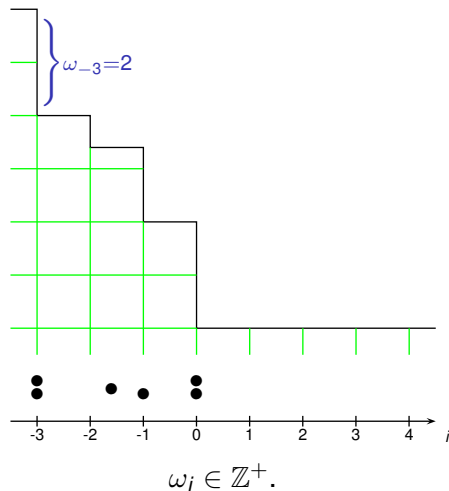
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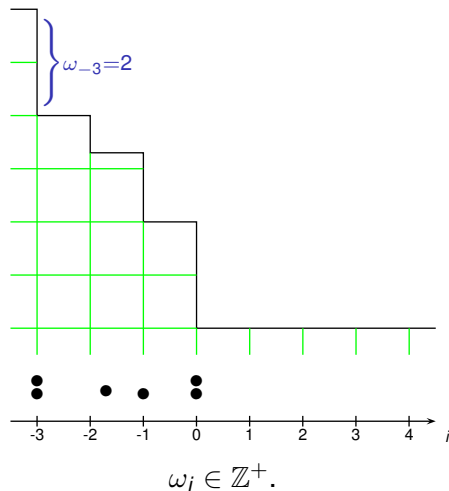
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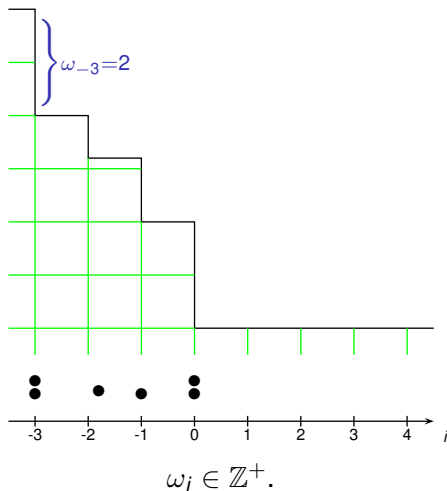
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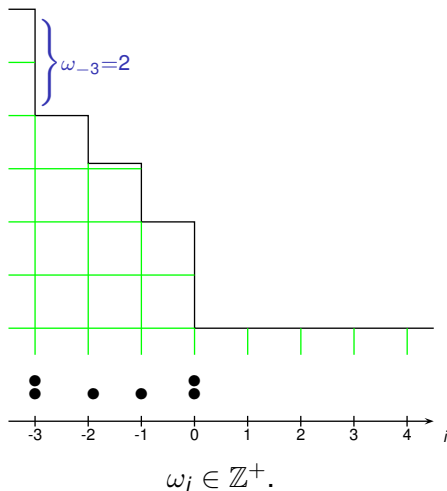
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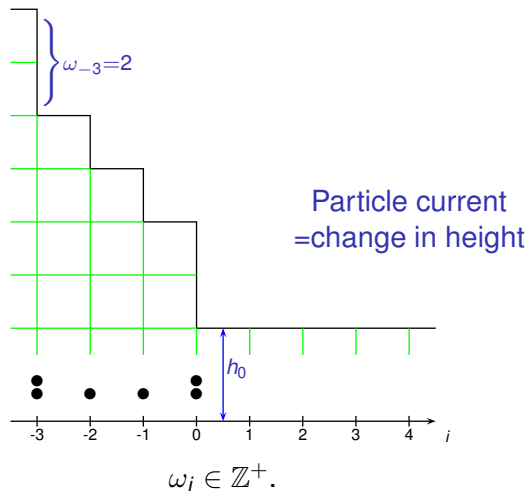
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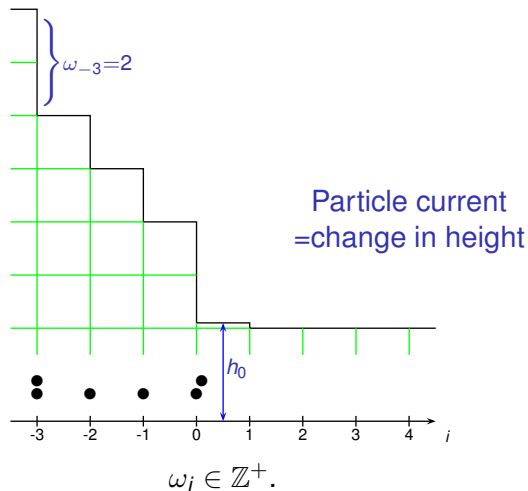
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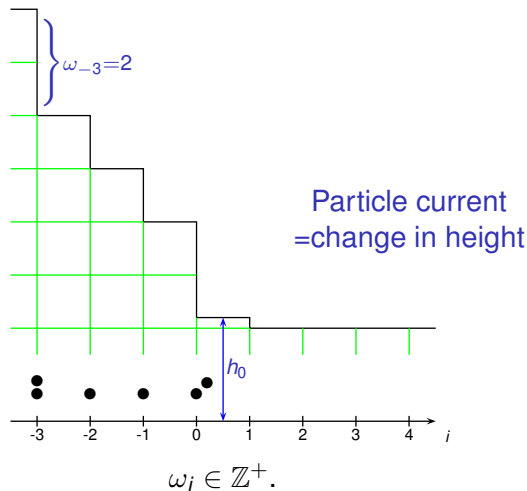
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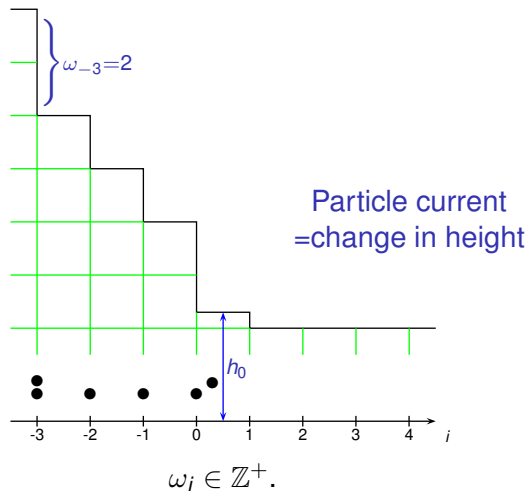
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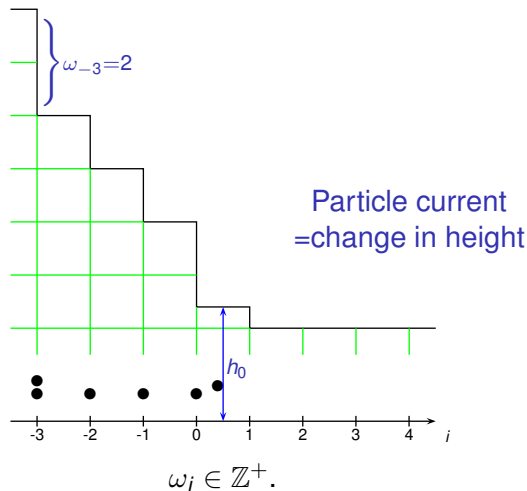
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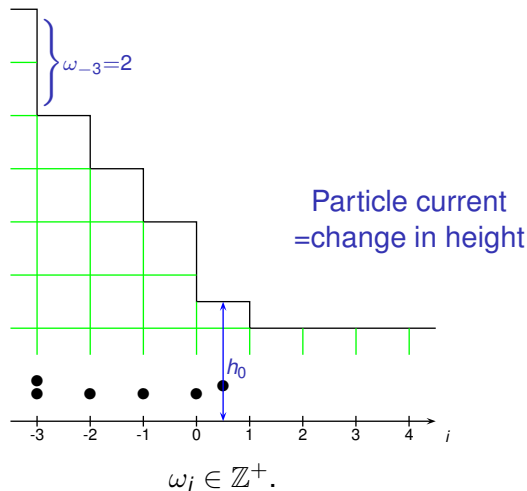
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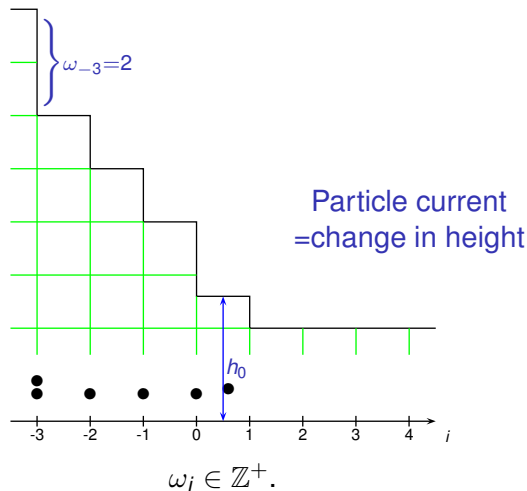
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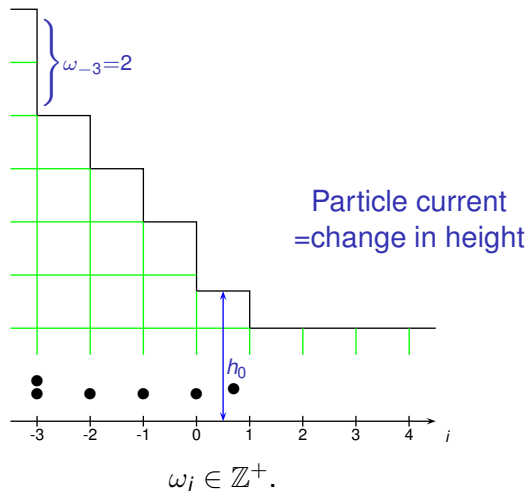
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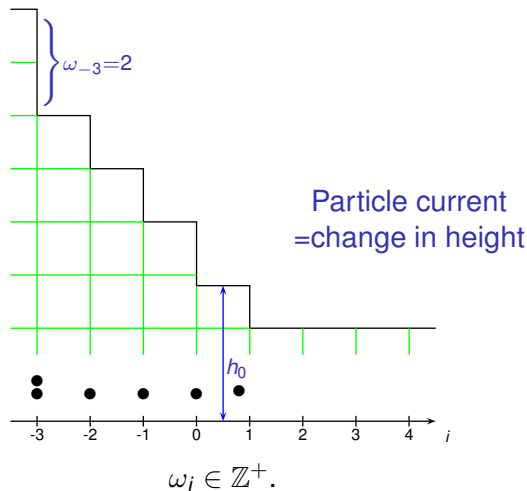
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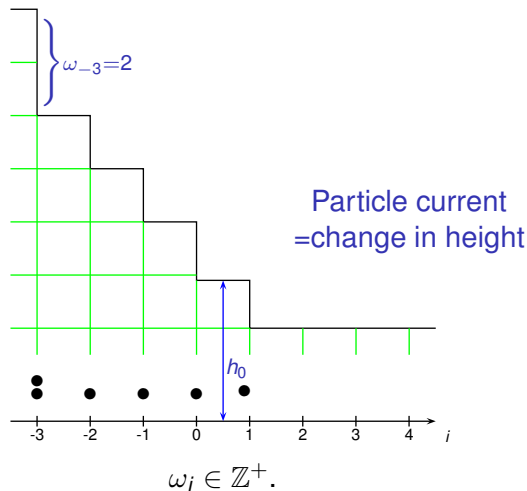
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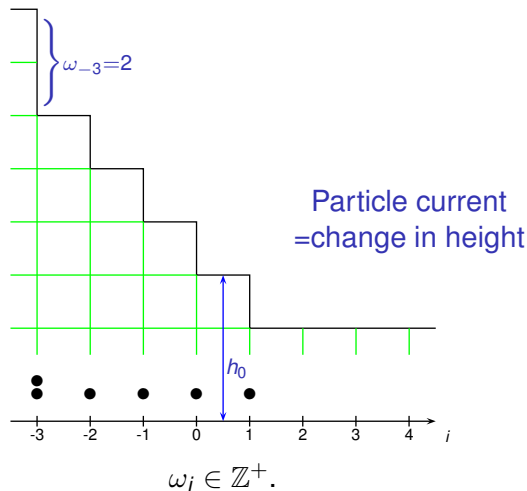
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The asymmetric zero range process

We need r non-decreasing.

Examples:

- ▶ 'Classical' ZRP: $r(\omega_i) = \mathbf{1}\{\omega_i > 0\}$.
- ▶ Independent walkers: $r(\omega_i) = \omega_i$.

Translation invariant measures

No surprise that constant density is stationary:

ASEP: $\eta_i \sim \text{iid. Bernoulli}(\rho)$.

Classical ZRP: $\omega_j \sim \text{iid. Geometric}(\frac{1}{1+\rho})$.

Independent walkers: $\omega_j \sim \text{iid. Poisson}(\rho)$.

These are the only extremal translation-invariant distributions.

Hydrodynamics: asymmetric ZRP

Take $p > q = 1 - p$, and AZRP with right rate $p \cdot r(\omega_i)$, left rate $q \cdot r(\omega_i)$.

$$\frac{d}{d\tau} \mathbf{E}\omega_i = p \mathbf{E}r(\omega_{i-1}) + q \mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_i)$$

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Ballistic scaling (zoom out and speed up by factor L):

- ▶ $\varrho(t, x) = \mathbf{E}\omega_{Lx}(Lt)$;
- ▶ also define $G(\varrho) = \mathbf{E}^{\varrho}r(\omega)$:

$$\frac{d}{d(\tau/L)} \mathbf{E}\omega_i = qL(\mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_i)) - pL(\mathbf{E}r(\omega_i) - \mathbf{E}r(\omega_{i-1}))$$

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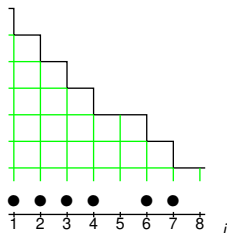
$$\begin{aligned} \frac{d}{d(\tau/L)} \mathbf{E}\omega_i &= qL(\mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_i)) - pL(\mathbf{E}r(\omega_i) - \mathbf{E}r(\omega_{i-1})) \\ \frac{\partial}{\partial t} \varrho(t, x) &= (q - p) \frac{\partial}{\partial x} G(\varrho(t, x)). \end{aligned}$$

Hydrodynamics: asymmetric ZRP

$$\frac{\partial}{\partial t} \varrho(t, x) + (p - q) \frac{\partial}{\partial x} G(\varrho(t, x)) = 0$$

Classical ZRP: $G(\varrho) = \mathbf{E}^{\varrho} r(\omega) = \mathbf{E}^{\varrho} \mathbf{1}\{\omega > 0\} = \frac{\varrho}{1+\varrho}$
 concave, **Burgers-type equation**.

Independent walkers: $G(\varrho) = \mathbf{E}^{\varrho} r(\omega) = \mathbf{E}^{\varrho} \omega = \varrho$
 linear, **transport equation**.



The stationary solution is constant density,
 linear slope.

Hydrodynamics: symmetric ZRP

Take $p = q = \frac{1}{2}$, and SZRP with right rate $\frac{1}{2}r(\omega_i)$, left rate $\frac{1}{2}r(\omega_i)$.

$$\frac{d}{d\tau} \mathbf{E}\omega_i = \frac{1}{2} \mathbf{E}r(\omega_{i-1}) + \frac{1}{2} \mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_i)$$

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Diffusive scaling:

- ▶ $\varrho(t, x) = \mathbf{E}\omega_{Lx}(L^2t)$;
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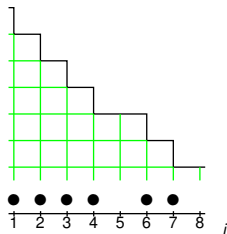
$$\frac{\partial}{\partial t} \varrho(t, x) = \frac{1}{2} \frac{\partial^2}{\partial x^2} G(\varrho(t, x)).$$

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 linear, **heat equation.**



The stationary solution is constant density, linear slope, or linearly changing G with current.

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Can we model sedimentation and erosion processes with these surfaces?

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Hills

Can we model sedimentation and erosion processes with these surfaces?

Issues:

- ▶ Hills are not always straight \leftrightarrow translation invariance.
- ▶ Most hillslopes are rather stationary \leftrightarrow particle current.

Convex hills



Wikipedia

Concave hills



Product blocking measures

Solution: block particles (**no current**) and make their rates asymmetric (**non-constant density**).

Can we have a reversible stationary distribution in product form:

$$\underline{\mu}(\underline{\omega}) = \bigotimes_i \mu_i(\omega_i);$$

$$\underline{\mu}(\underline{\omega}) \cdot \text{rate}(\underline{\omega} \rightarrow \underline{\omega}^{i \curvearrowright i+1}) = \underline{\mu}(\underline{\omega}^{i \curvearrowright i+1}) \cdot \text{rate}(\underline{\omega}^{i \curvearrowright i+1} \rightarrow \underline{\omega}) \quad ?$$

Here

$$\underline{\omega}^{i \curvearrowright i+1} = \underline{\omega} - \underline{\delta}_i + \underline{\delta}_{i+1}.$$

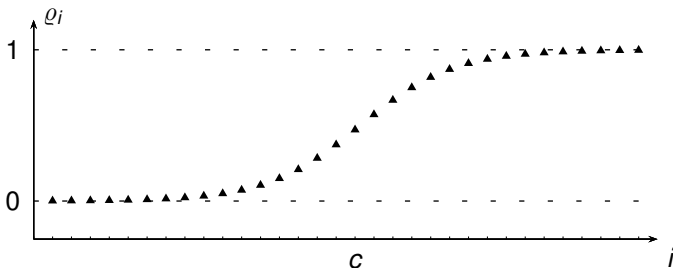
Asymmetric simple exclusion

$$\underline{\mu}(\underline{\eta}) \cdot \text{rate}(\underline{\eta} \rightarrow \underline{\eta}^{i \leftrightarrow i+1}) = \underline{\mu}(\underline{\eta}^{i \leftrightarrow i+1}) \cdot \text{rate}(\underline{\eta}^{i \leftrightarrow i+1} \rightarrow \underline{\eta})$$

ASEP: $\mu_i \sim \text{Bernoulli}(\varrho_i)$; $\bullet \longrightarrow \eta$

$$\varrho_i(1 - \varrho_{i+1}) \cdot p = (1 - \varrho_i)\varrho_{i+1} \cdot q$$

Solution:
$$\varrho_i = \frac{(\frac{p}{q})^{i-c}}{1 + (\frac{p}{q})^{i-c}} = \frac{1}{(\frac{q}{p})^{i-c} + 1}$$



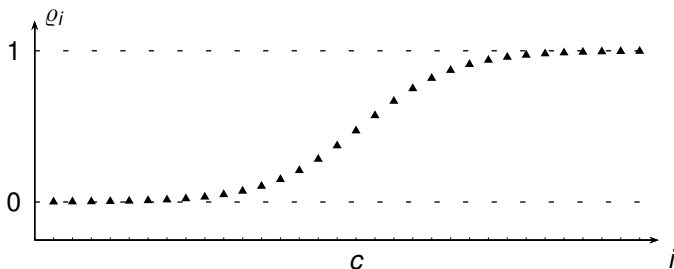
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Asymmetric zero range process

$$\underline{\mu}(\underline{\omega}) \cdot \text{rate}(\underline{\omega} \rightarrow \underline{\omega}^{i \curvearrowright i+1}) = \underline{\mu}(\underline{\omega}^{i \curvearrowright i+1}) \cdot \text{rate}(\underline{\omega}^{i \curvearrowright i+1} \rightarrow \underline{\omega}) \quad ?$$

AZRP, classical:

$$\mu_i(\omega_i) \mu_{i+1}(\omega_{i+1}) \cdot p \mathbf{1}\{\omega_i > 0\} = \mu_i(\omega_i - 1) \mu_{i+1}(\omega_{i+1} + 1) \cdot q$$

Solution: $\mu_i \sim \text{Geometric}\left(1 - \left(\frac{p}{q}\right)^{i - \text{const}}\right).$

Asymmetric zero range process

$$\underline{\mu}(\underline{\omega}) \cdot \text{rate}(\underline{\omega} \rightarrow \underline{\omega}^{i \leftrightarrow i+1}) = \underline{\mu}(\underline{\omega}^{i \leftrightarrow i+1}) \cdot \text{rate}(\underline{\omega}^{i \leftrightarrow i+1} \rightarrow \underline{\omega}) \quad ?$$

AZRP, independent walkers:

$$\mu_i(\omega_i) \mu_{i+1}(\omega_{i+1}) \cdot p \omega_i = \mu_i(\omega_i - 1) \mu_{i+1}(\omega_{i+1} + 1) \cdot q(\omega_{i+1} + 1)$$

Solution: $\mu_i \sim \text{Poisson}\left(\left(\frac{p}{q}\right)^{i-\text{const}}\right).$

Further models

In fact product blocking measures are very general.

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- ▶ q -exclusion
- ▶ Katz-Lebowitz-Spohn model

Product blocking measures

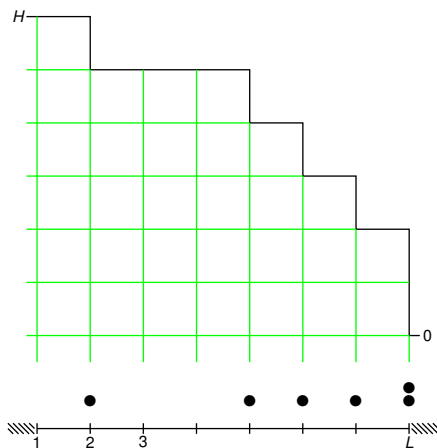
They are also very handy, due to reversibility.

Take a stationary, reversible Markov chain. Cut any of its edges. It stays reversible stationary w.r.t. the same distribution.

In our case: freeze the boundaries to obtain a stationary hill slope.

Microscopic model

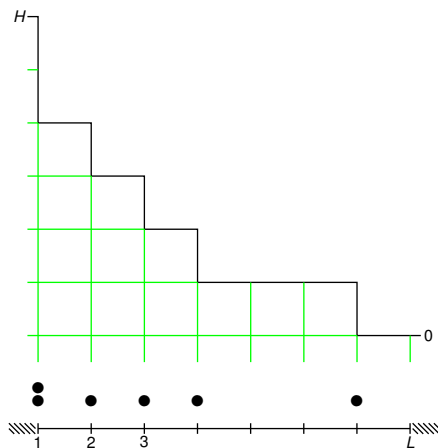
Our choice: AZRP with frozen boundaries. $p > q$: convex



Particles jump to the right with rate $p \cdot r(\omega_i)$
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Particles jump to the right with rate $p \cdot r(\omega_i)$
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Microscopic model

Notice:

- ▶ The height of the hill H is conserved, **the product measure is *not* ergodic.**
- ▶ One-site marginals, given H , are in general not explicit.
- ▶ Except for independent walkers, where ω_j are Binomial.

Microscopic model

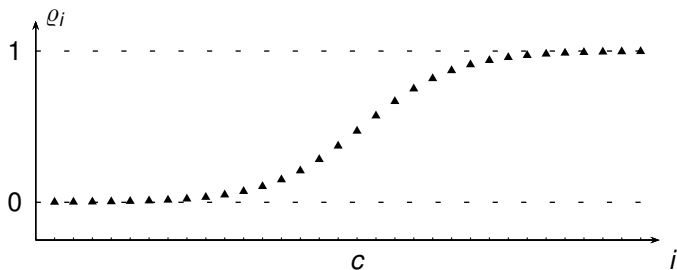
Notice:

- ▶ The height of the hill H is conserved, **the product measure is *not* ergodic.**
- ▶ One-site marginals, given H , are in general not explicit.
- ▶ Except for independent walkers, where ω_j are Binomial.

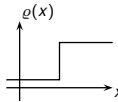
We won't be bothered by this.

Hydrodynamics

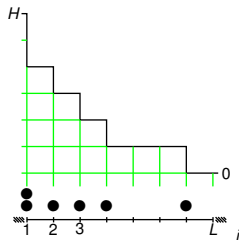
Work in progress...



A blocking measure is a microscopic object. Here is its scaling

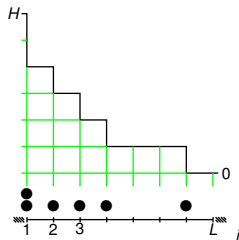
limit:  , not very interesting.

Hydrodynamics



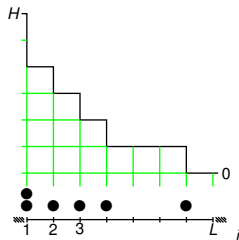
► Scaling parameter: L

Hydrodynamics



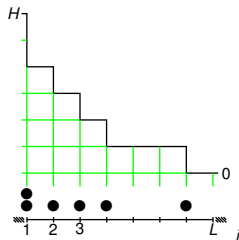
- ▶ Scaling parameter: L
- ▶ Blocking measure marginals depend on $\left(\frac{p}{q}\right)^i = \left(\frac{p}{q}\right)^{Lx}$.

Hydrodynamics



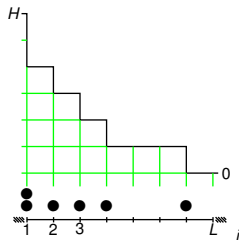
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Hydrodynamics



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- ▶ Then check $\frac{d}{d\tau} \mathbf{E}\omega_i(\tau)$.

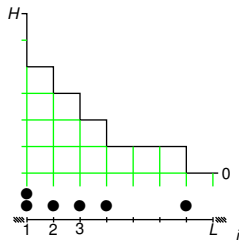
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Hydrodynamics



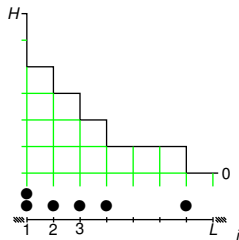
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For AZRP (rates $p \cdot r(\omega_i)$ right and $q \cdot r(\omega_i)$ left):

$$\begin{aligned} \frac{d}{d\tau} \mathbf{E}\omega_i &= \frac{1}{2} (\mathbf{E}r(\omega_{i-1}) - 2\mathbf{E}r(\omega_i) + \mathbf{E}r(\omega_{i+1})) \\ &\quad - \frac{\gamma}{L} (\mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_{i-1})). \end{aligned}$$

Hydrodynamics



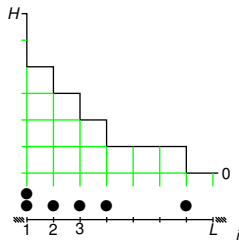
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Hydrodynamics

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which dictates **diffusive scaling**:

- ▶ $p = \frac{1}{2} + \frac{\gamma}{L}$, $q = \frac{1}{2} - \frac{\gamma}{L}$;
- ▶ $\varrho(t, x) = \mathbf{E}\omega_{Lx}(L^2t)$;
- ▶ also define $G(\varrho) = \mathbf{E}e^{\varrho r(\omega)}$:

$$\frac{d}{d(\tau/L^2)} \mathbf{E}\omega_i = \frac{L^2}{2} (\mathbf{E}r(\omega_{i-1}) - 2\mathbf{E}r(\omega_i) + \mathbf{E}r(\omega_{i+1})) - \gamma L (\mathbf{E}r(\omega_{i+1}) - \mathbf{E}r(\omega_{i-1})),$$

$$\frac{\partial}{\partial t} \varrho(t, x) = \frac{1}{2} \frac{\partial^2}{\partial x^2} G(\varrho(t, x)) - 2\gamma \frac{\partial}{\partial x} G(\varrho(t, x)), \quad (0 < x < 1).$$

Hydrodynamics

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How about the boundaries?

Hydrodynamics

- ▶ $\mathbf{p} = \frac{1}{2} + \frac{\gamma}{L}$, $\mathbf{q} = \frac{1}{2} - \frac{\gamma}{L}$;
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How about the boundaries?

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$$\frac{1}{L} \frac{d}{d(\tau/L^2)} \mathbf{E}\omega_1 = \frac{L}{2} (\mathbf{E}r(\omega_2) - \mathbf{E}r(\omega_1)) - \gamma (\mathbf{E}r(\omega_2) + \mathbf{E}r(\omega_1)),$$

Hydrodynamics

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$$0 = \frac{1}{2} \frac{\partial}{\partial x} G(\varrho(t, 0)) - 2\gamma G(\varrho(t, 0)).$$

Hydrodynamics

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Also: $0 = \frac{1}{2} \frac{\partial}{\partial x} G(\varrho(t, 1)) - 2\gamma G(\varrho(t, 1)).$

Hydrodynamics

$$\frac{\partial}{\partial t} \varrho(t, x) = \frac{1}{2} \frac{\partial^2}{\partial x^2} G(\varrho(t, x)) - 2\gamma \frac{\partial}{\partial x} G(\varrho(t, x)) \quad (0 < x < 1)$$

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$$0 = \frac{1}{2} \frac{\partial}{\partial x} G(\varrho(t, 1)) - 2\gamma G(\varrho(t, 1))$$

Convection-diffusion type equation with Robin boundary.

Hydrodynamics

$$\begin{aligned}\frac{\partial}{\partial t} \varrho(t, x) &= \frac{1}{2} \frac{\partial^2}{\partial x^2} G(\varrho(t, x)) - 2\gamma \frac{\partial}{\partial x} G(\varrho(t, x)) \quad (0 < x < 1) \\ &= \frac{\partial}{\partial x} \left(\frac{1}{2} \frac{\partial}{\partial x} G(\varrho(t, x)) - 2\gamma G(\varrho(t, x)) \right) \\ 0 &= \frac{1}{2} \frac{\partial}{\partial x} G(\varrho(t, 0)) - 2\gamma G(\varrho(t, 0)) \\ 0 &= \frac{1}{2} \frac{\partial}{\partial x} G(\varrho(t, 1)) - 2\gamma G(\varrho(t, 1))\end{aligned}$$

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Hydrodynamics

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 \end{aligned}$$

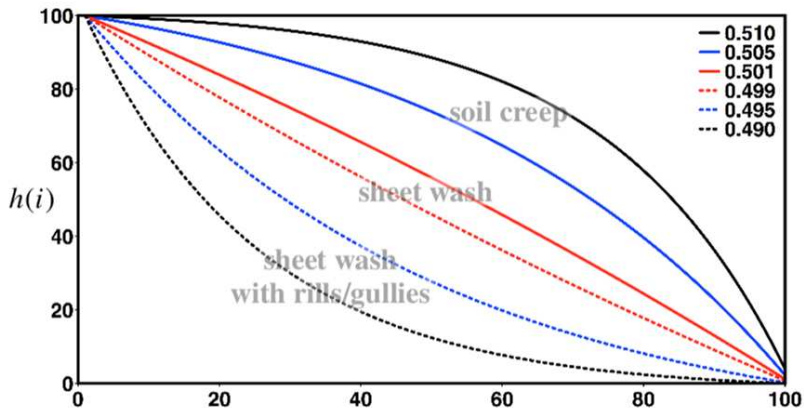
Convection-diffusion type equation with Robin boundary.

Doing the proper derivation is work in progress.

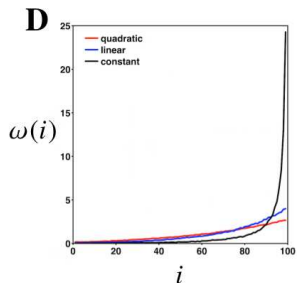
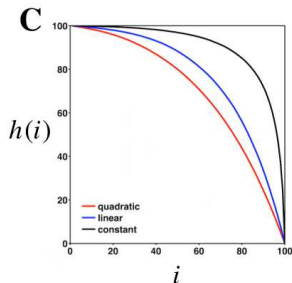
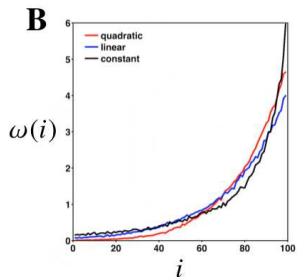
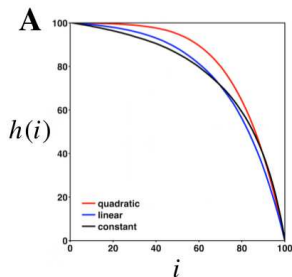
The time-stationary solution $G(\varrho(x)) = Ce^{4\gamma x}$ is consistent with the stationary blocking measure.

The stationary slope

$$G(\rho(x)) = Ce^{4\gamma x}$$



The stationary slope



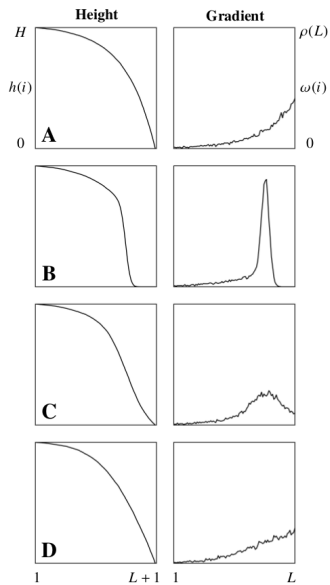
Dynamics

Space scale: $x \in [0, 1] \Leftrightarrow \text{we} \in \text{hill}$.

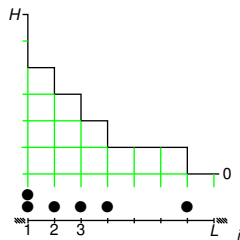
Problem 1: The stationary hillslope will not tell us the time scale.

\rightsquigarrow Observe relaxation to stationarity in Nature and in the PDE.

Dynamics



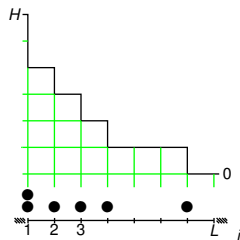
Dynamics



Problem 2: Geologists want a prediction for the *hill particle flux*, and the *distance travelled by hill particles*.

Notice: **Hill particles** \neq **our particles**.

Dynamics

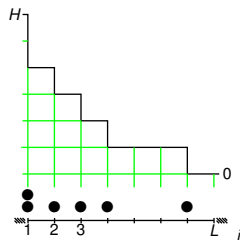


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Dynamics



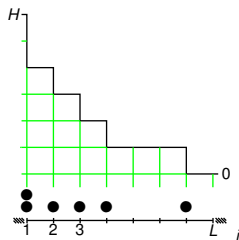
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Dynamics



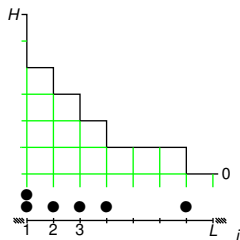
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Dynamics



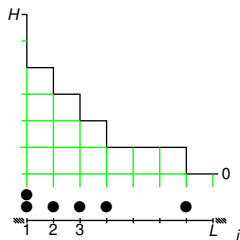
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- ▶ Average hill particle flux is the same across the hill (**reversibility**), *but this is not provided by the model*.

Dynamics



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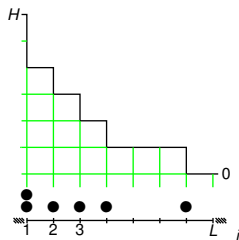
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One can then give an expected distance travelled by a hill particle.

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Thank you.